

Understanding drivers of geohazards using geochemical methods

(Jaime D. Barnes, University of Texas at Austin)

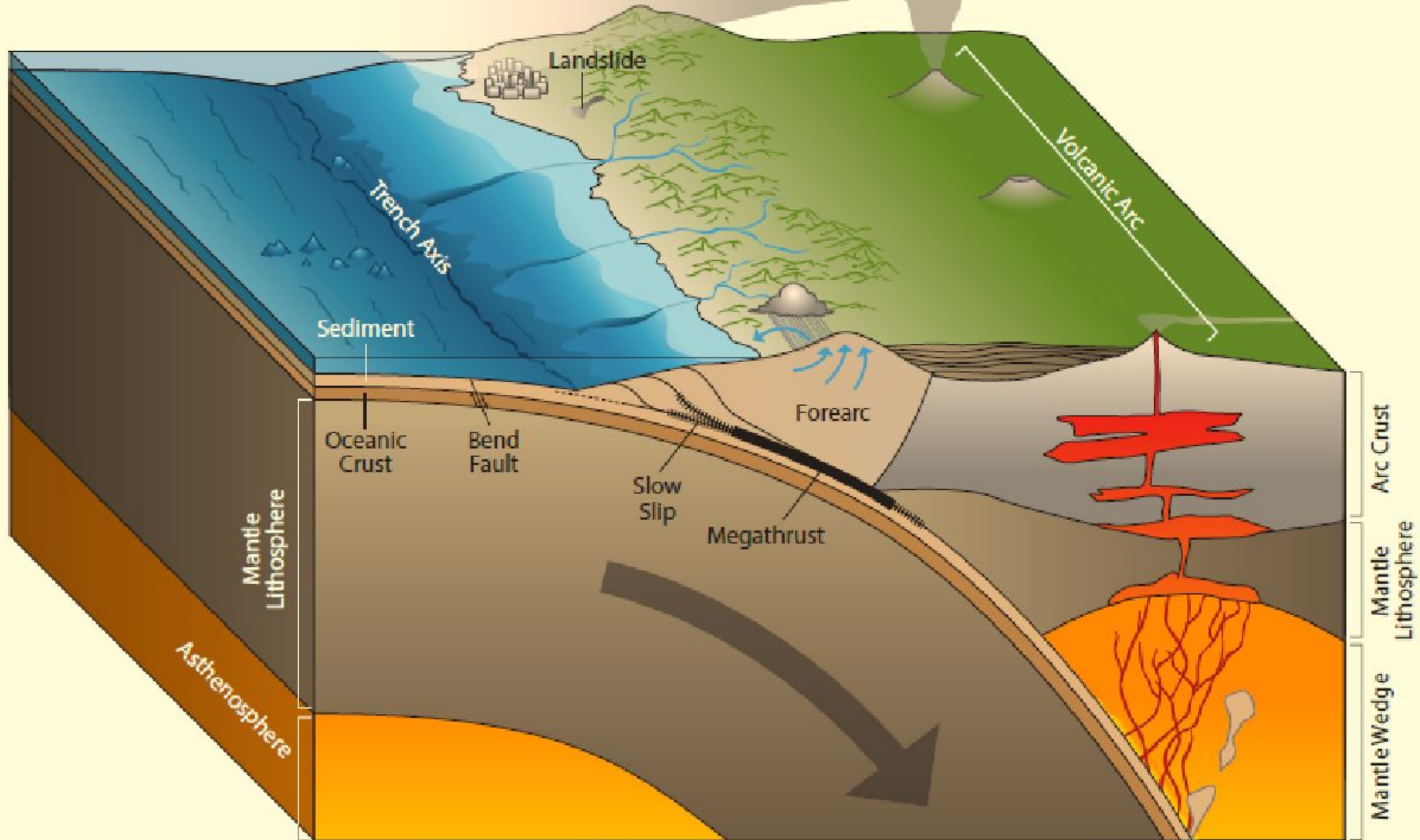
Fluids and fluid migration

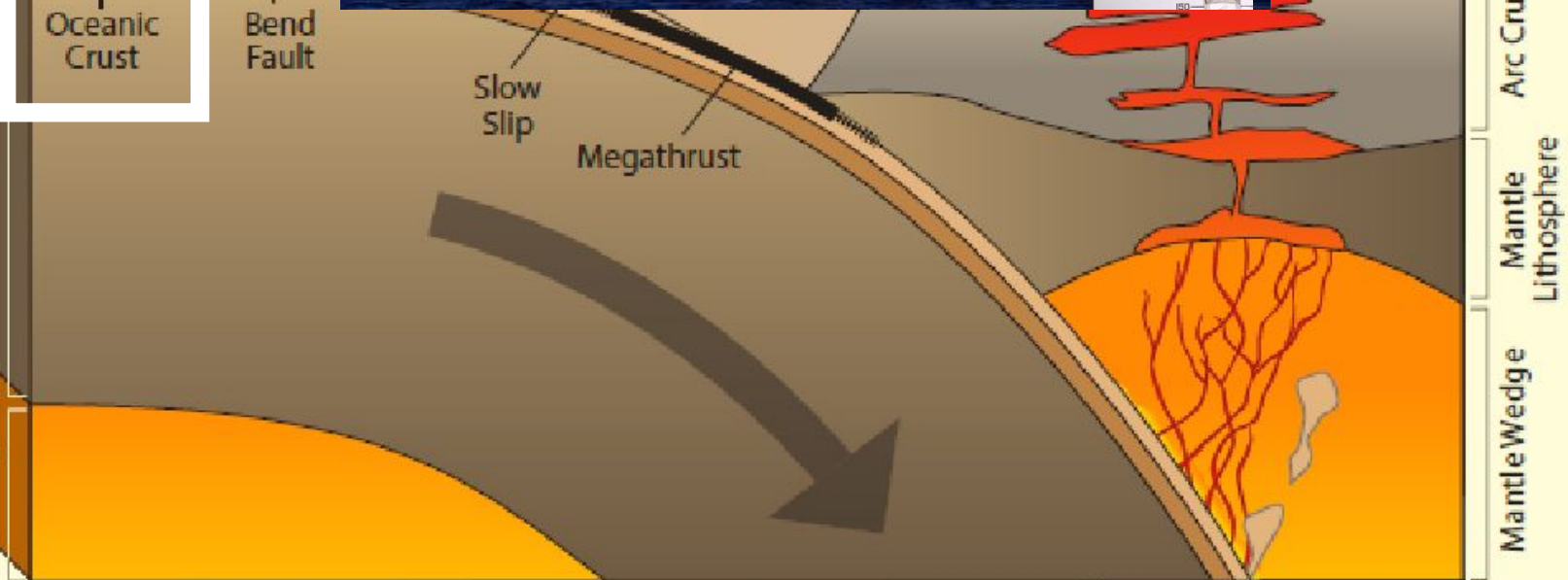
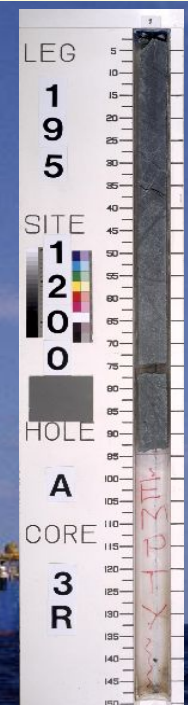
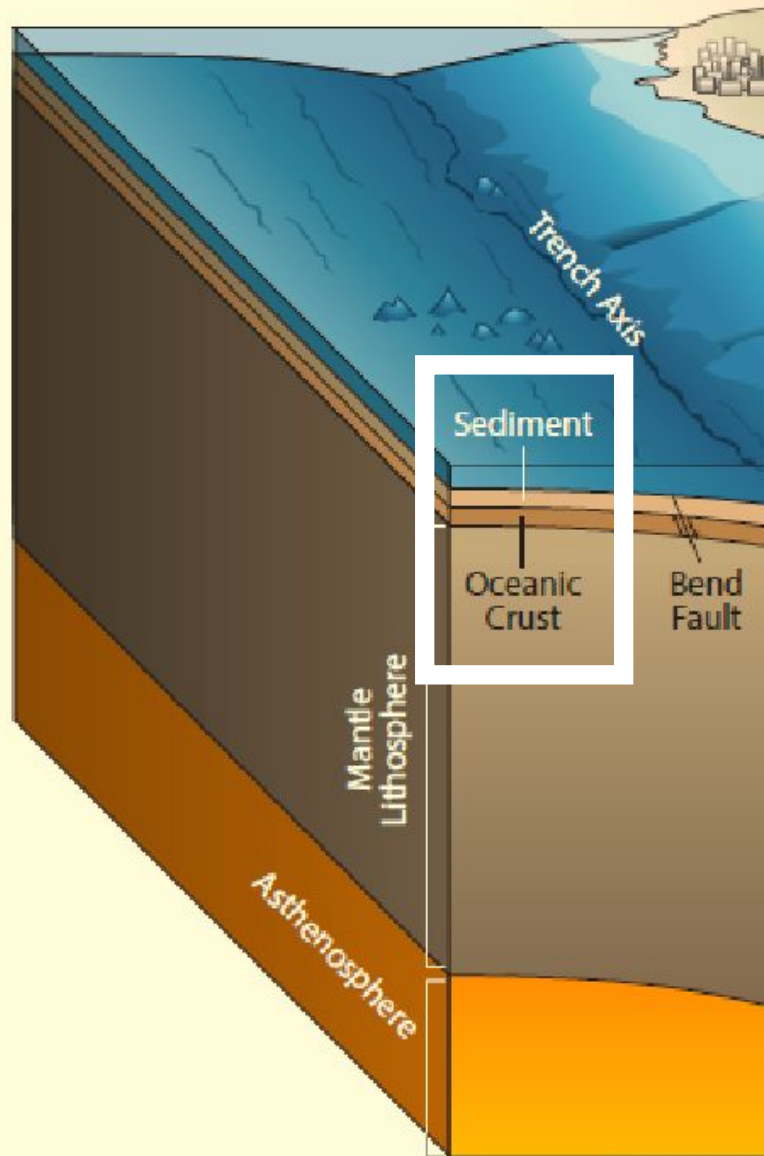
*How does fluid migration influence hazards and
material transport across the entire subduction
system?*

Fluids play a critical role in subduction zones:

- 1) Nature of seismic behavior along the megathrust (& associated faults)
- 2) Rock rheology and strain localization
- 3) Volcanic eruptive behavior
- 4) Resources: geothermal to critical metals

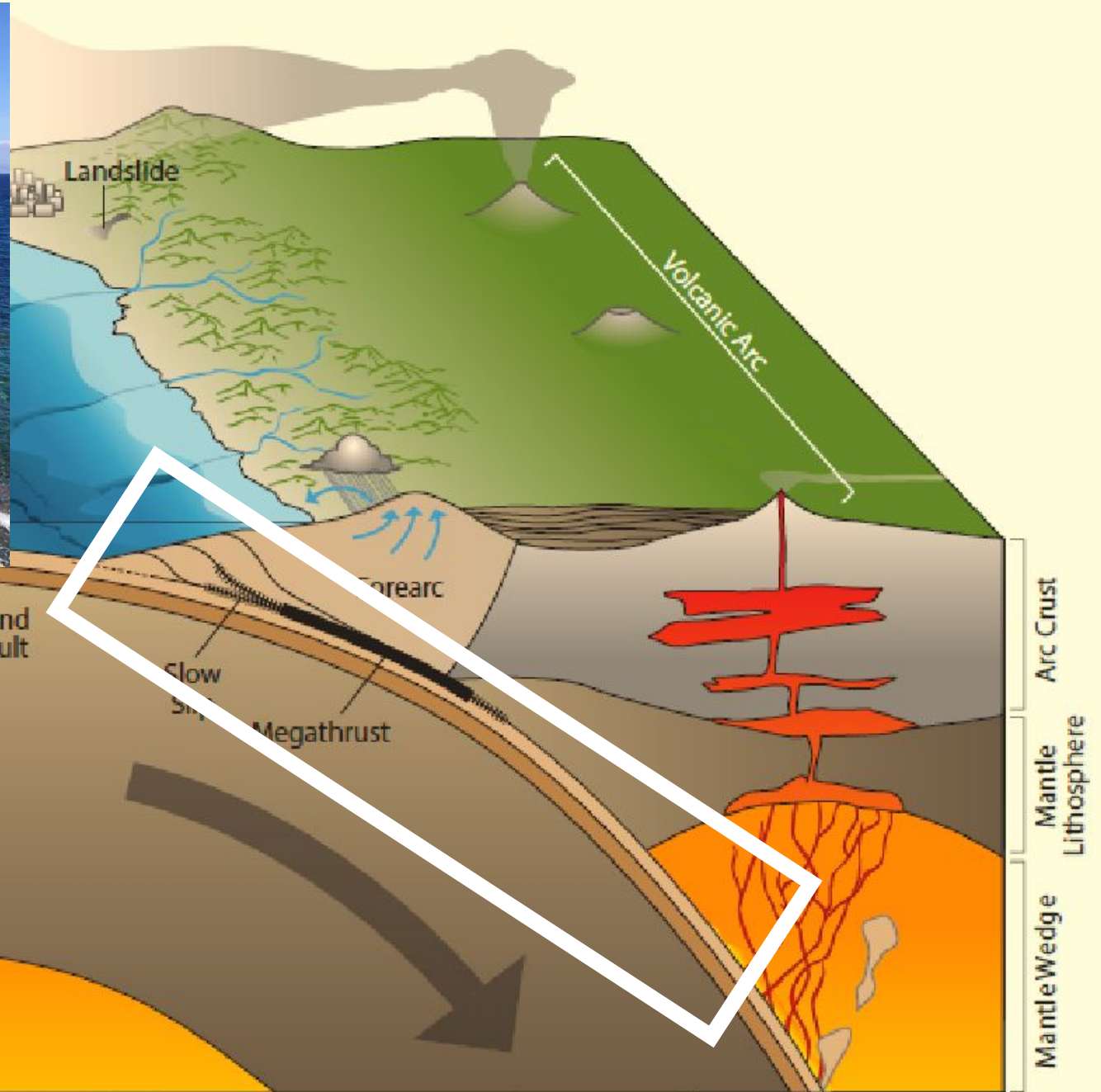
□ **fluids/volatiles/FME – amount? source? pathway?**
elemental concentrations, ratio, isotopic composition





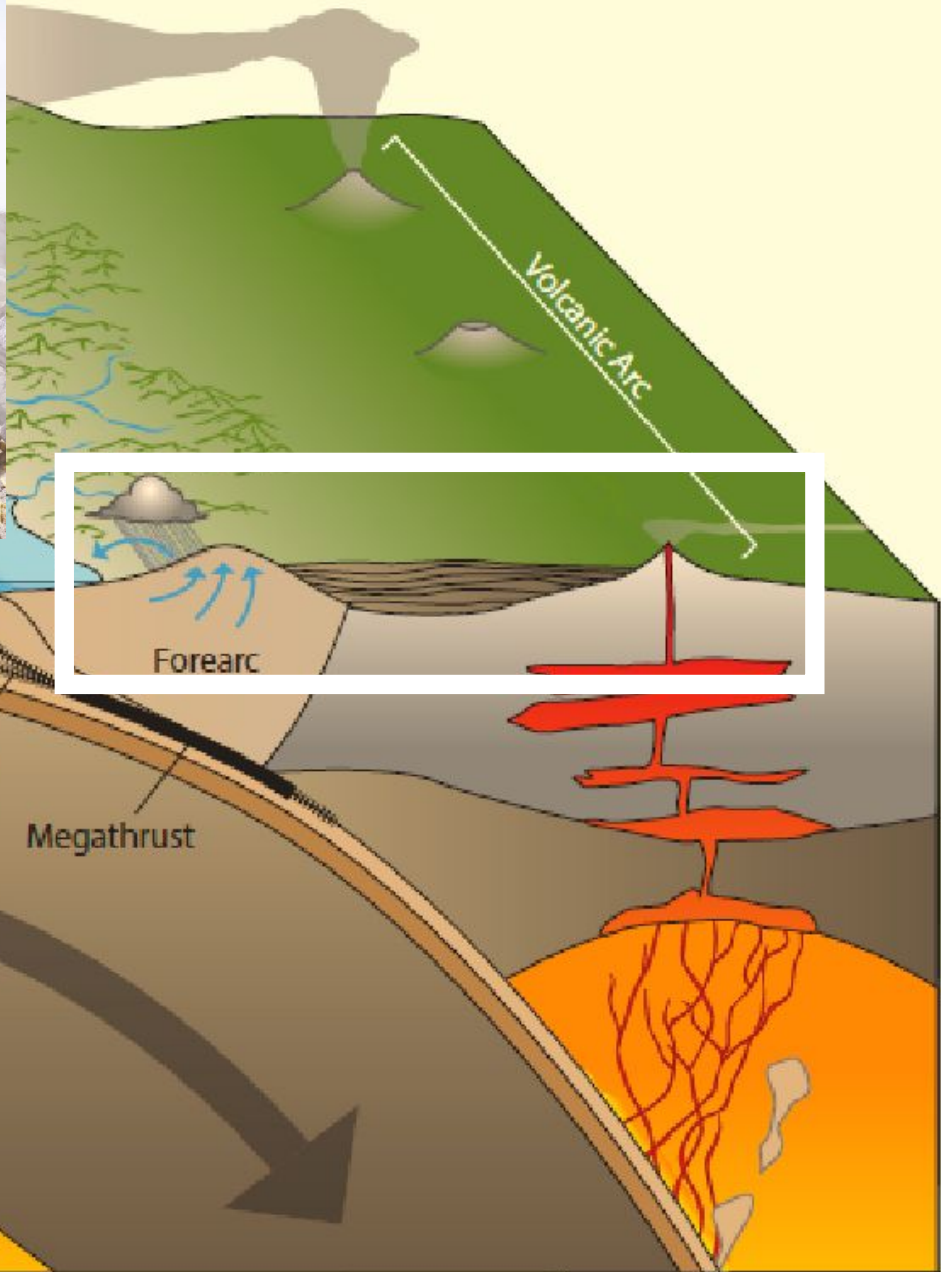
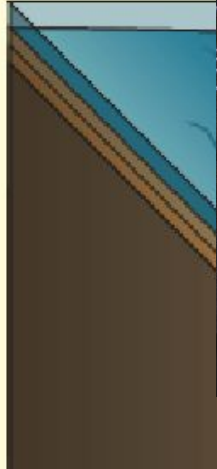


M. Cisneros (Syros)





M de Moor (Chile)
 photograph for *Nature* by J. Pasotti (June 2022)



Mangapakeha, New Zealand



Stable Isotope Geochemistry

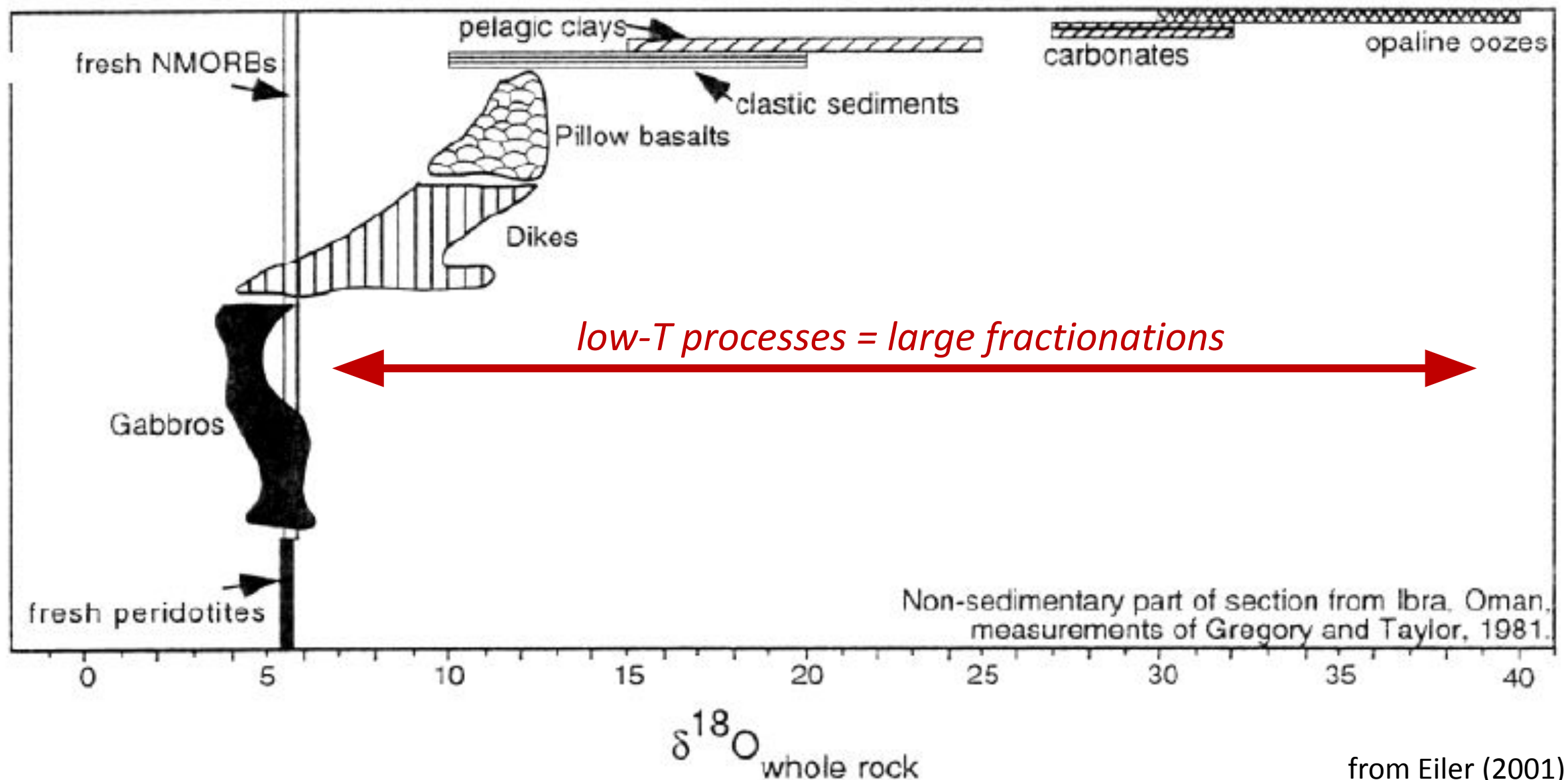
$$\delta^7\text{Li} = \left(\frac{{}^7\text{Li}/{}^6\text{Li}_{\text{sample}}}{{}^7\text{Li}/{}^6\text{Li}_{\text{standard}}} - 1 \right) \times 1000$$

□ measured on a mass spectrometer

Traditional: O, H, C, N, S

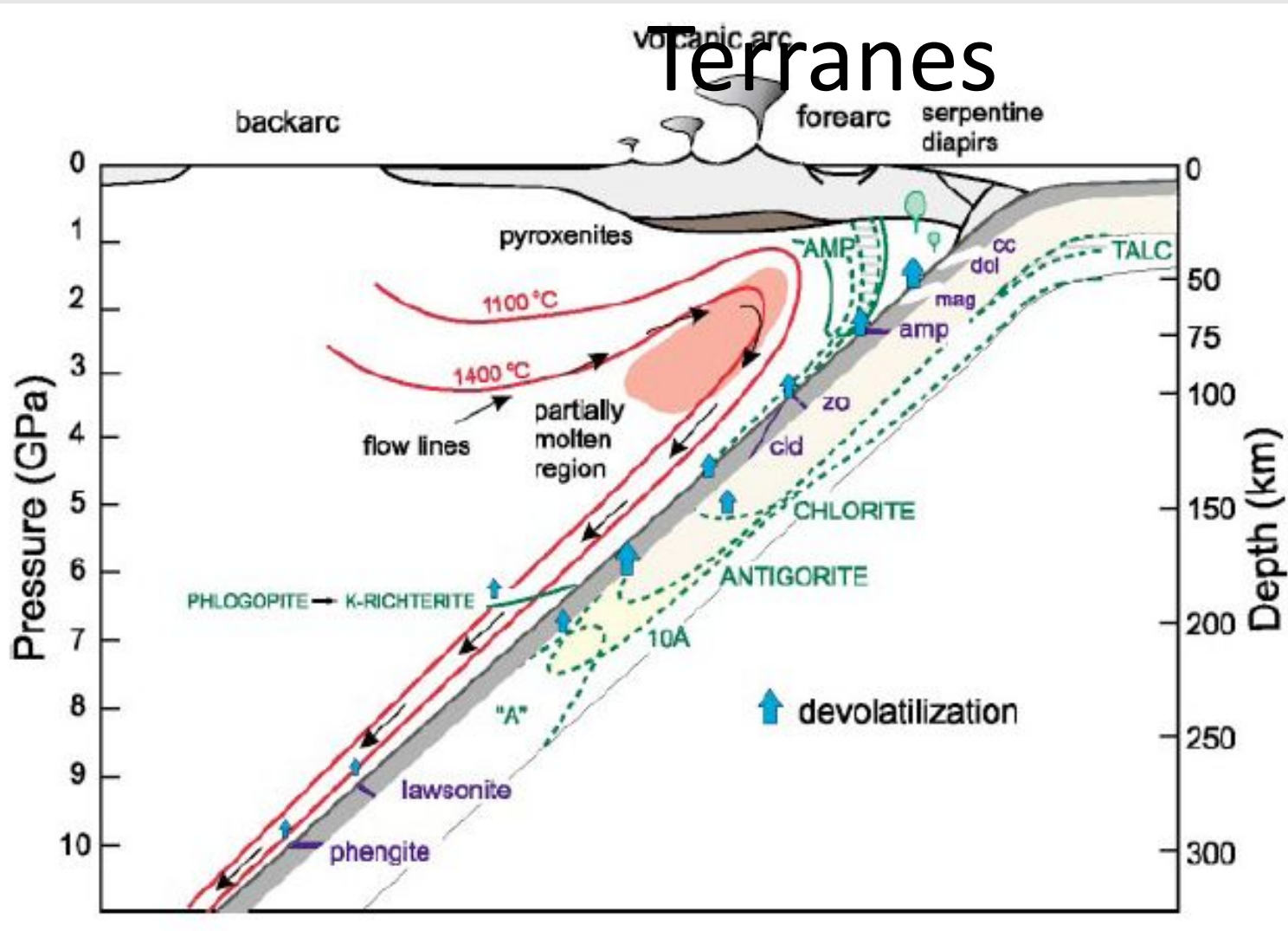
Non-traditional: Cl, Li, B, Ca, Fe, Cu, Zn...

Inputs



from Eiler (2001)

Exhumed Terranes



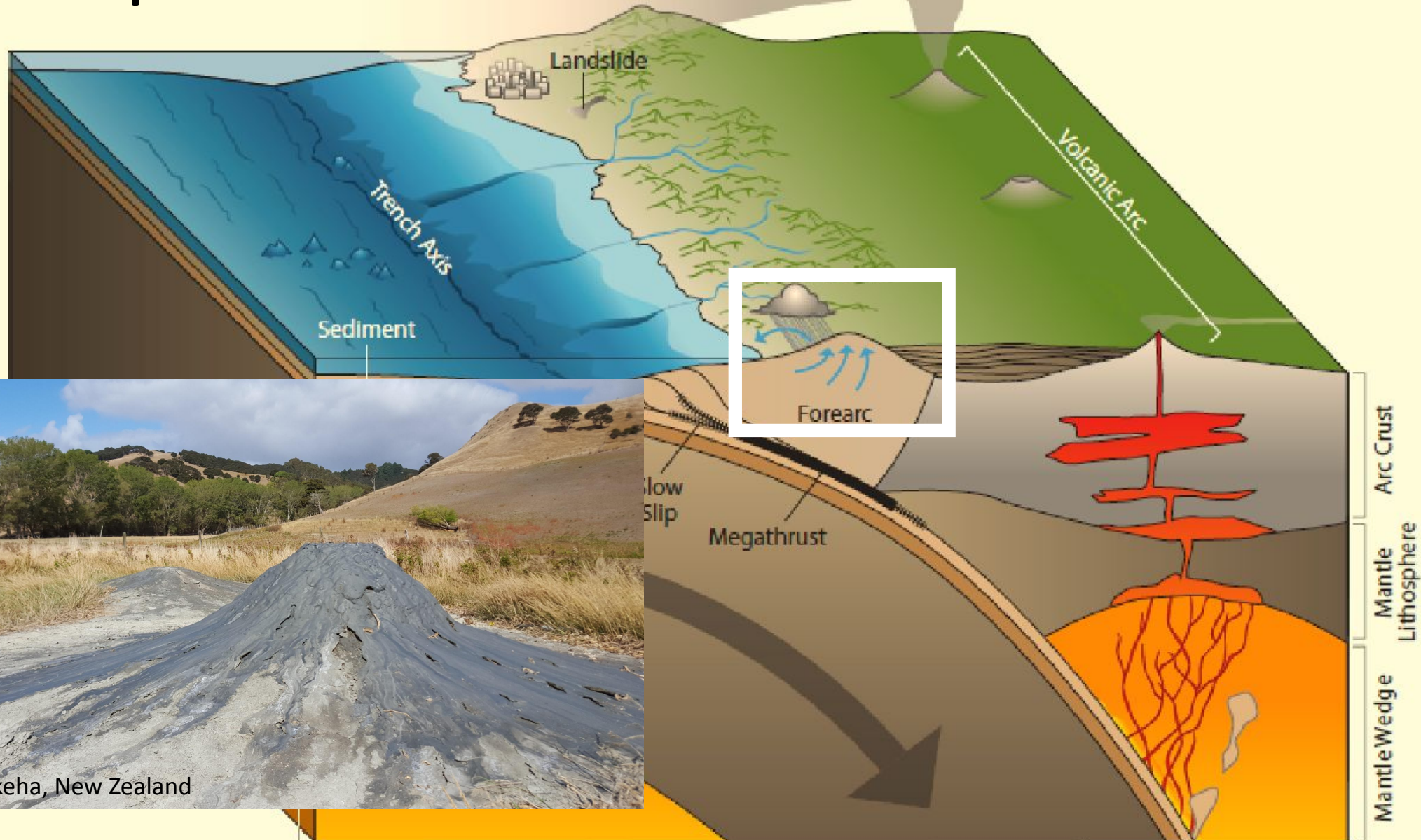
(Poli and Schmidt, 2002)

What is being lost where?

- concentrations
- elemental ratios
- isotopic compositions

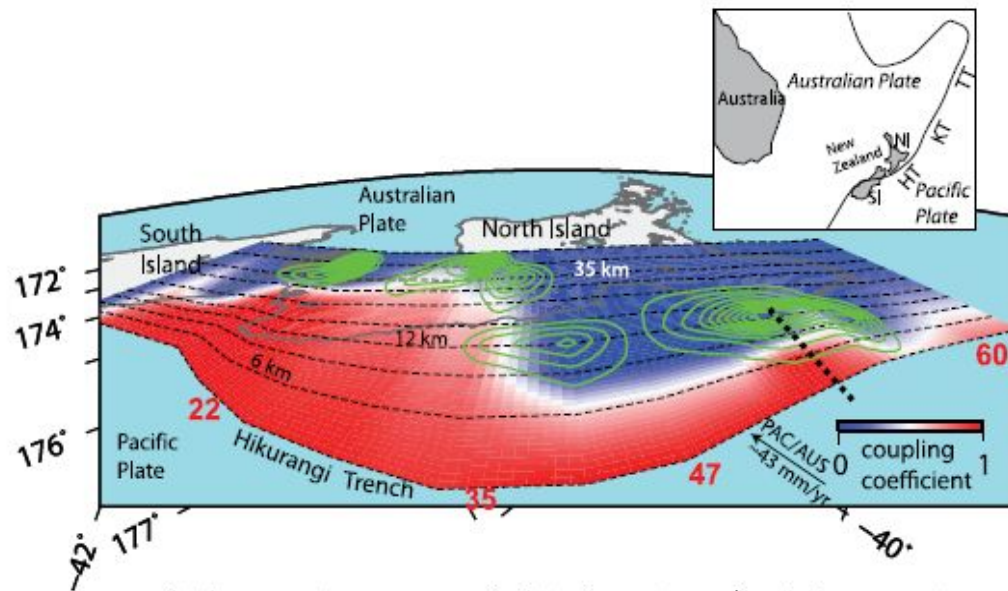
@ high-T fractionations
are small □ trace source

Outputs



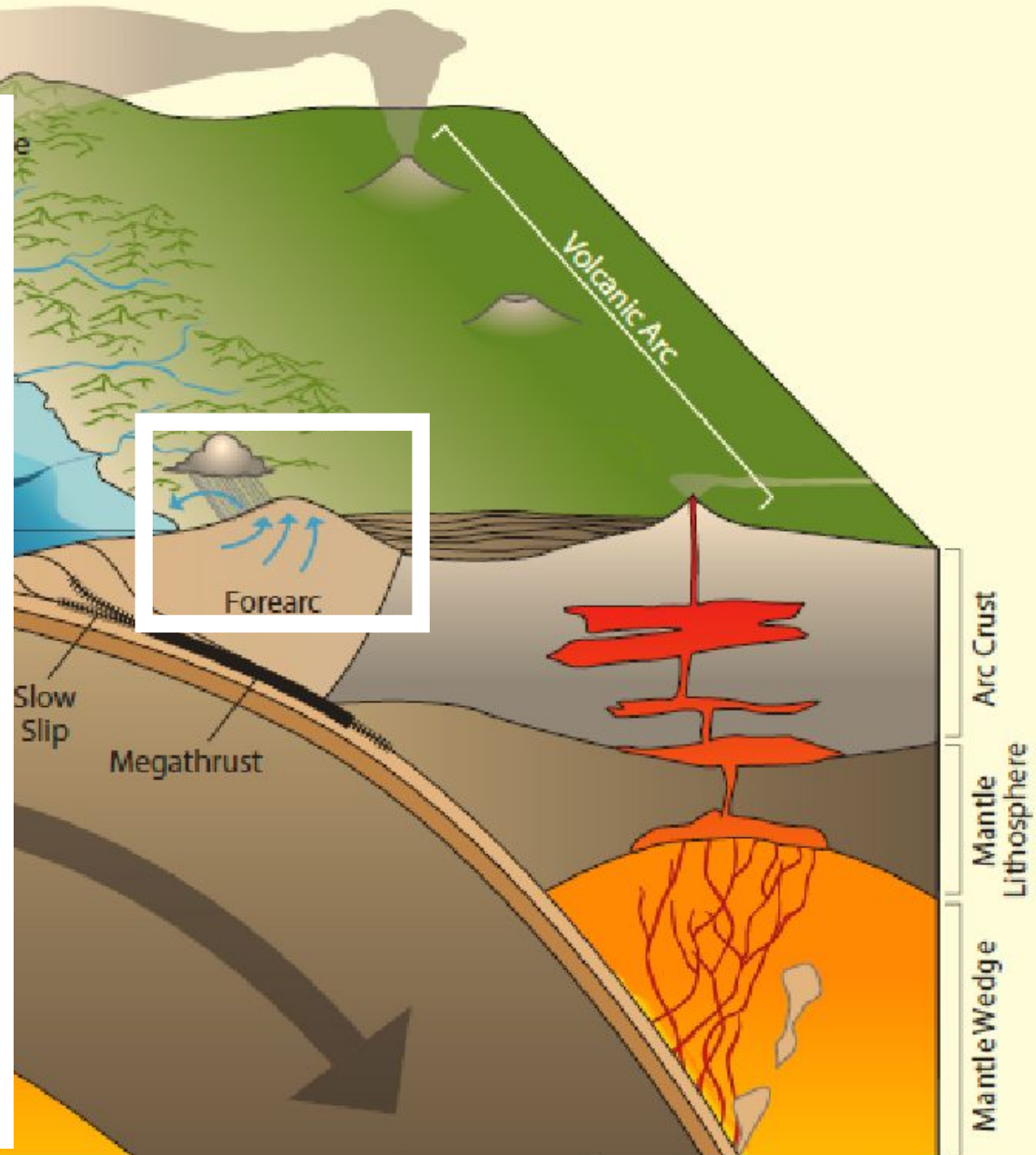
Mangapakeha, New Zealand

Outputs

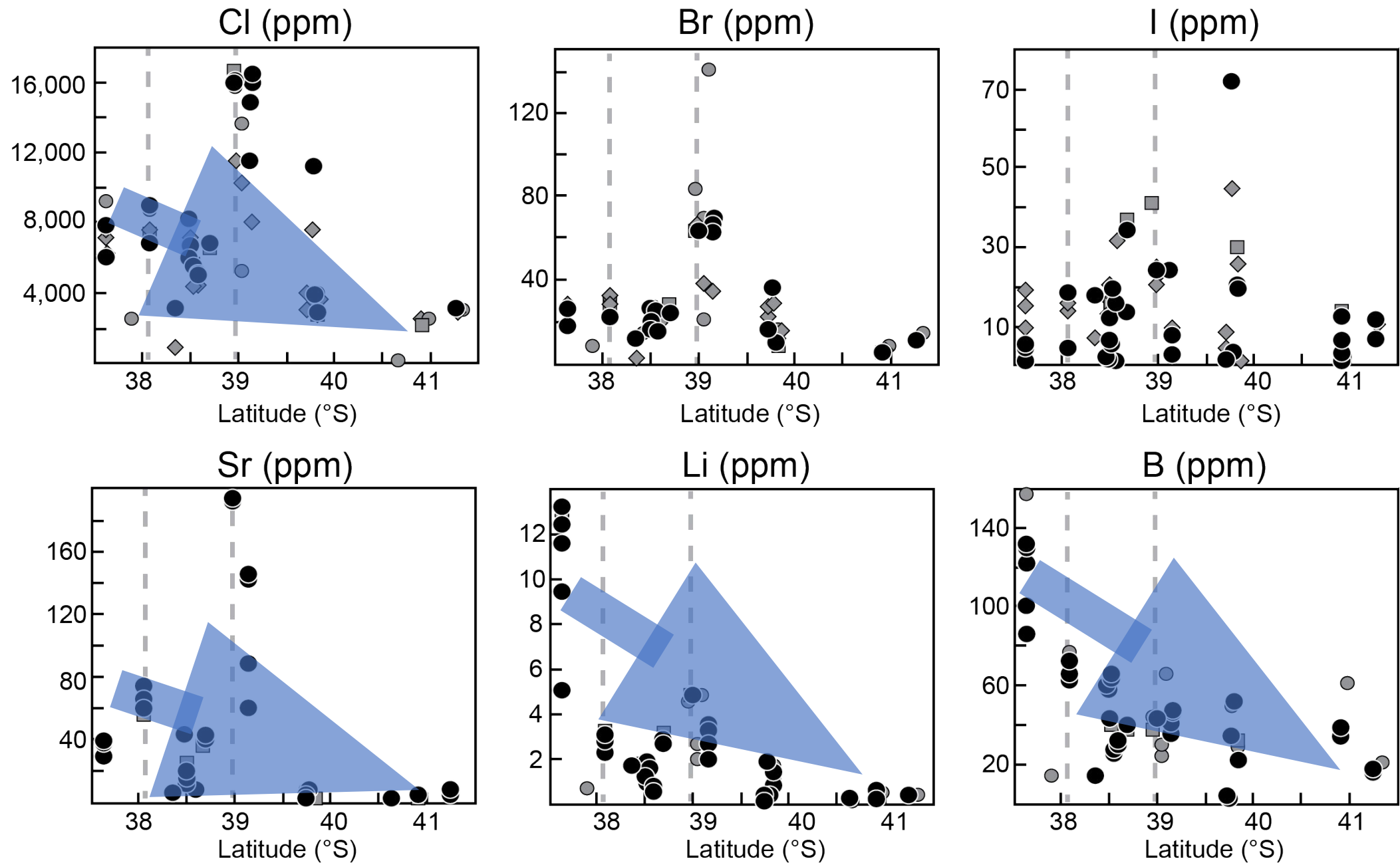


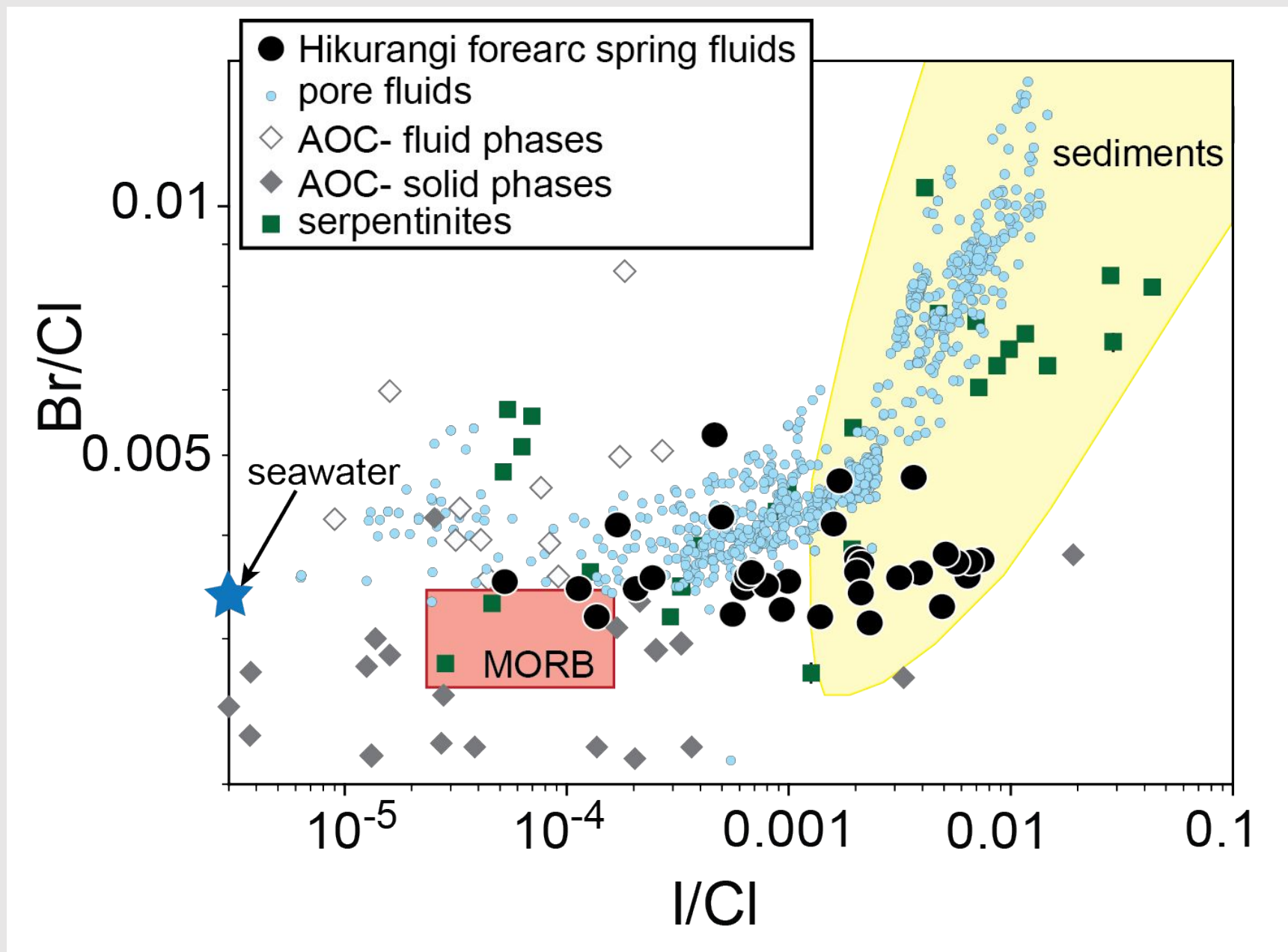
Southern segment (Wairarapa)	Central segment (Hawke's Bay)	Northern segment (Raukumara)
<ul style="list-style-type: none"> deep slow slip events (SSEs) and strong interseismic coupling frontal accretion dominant few seamounts entering margin 	<ul style="list-style-type: none"> shallow SSEs and weak interseismic coupling tectonic erosion/moderate-low accretion numerous seamounts impacting margin 	<ul style="list-style-type: none"> tectonic erosion/negligible accretion
← increasing thickness of sediment on the incoming plate		
<ul style="list-style-type: none"> tectonic contraction and strike-slip in upper plate 	<ul style="list-style-type: none"> tomographic evidence for high fluid content within upper plate back-arc extension and strike-slip in upper plate 	
← increasing convergence rate at trench		
<ul style="list-style-type: none"> wedge taper angle 4-6 degrees 	<ul style="list-style-type: none"> accretionary wedge taper angle 6-10 degrees 	
← decreasing Cl, B, Br, Na, Sr concentrations in fore-arc springs		

Wallace et al. (2010)

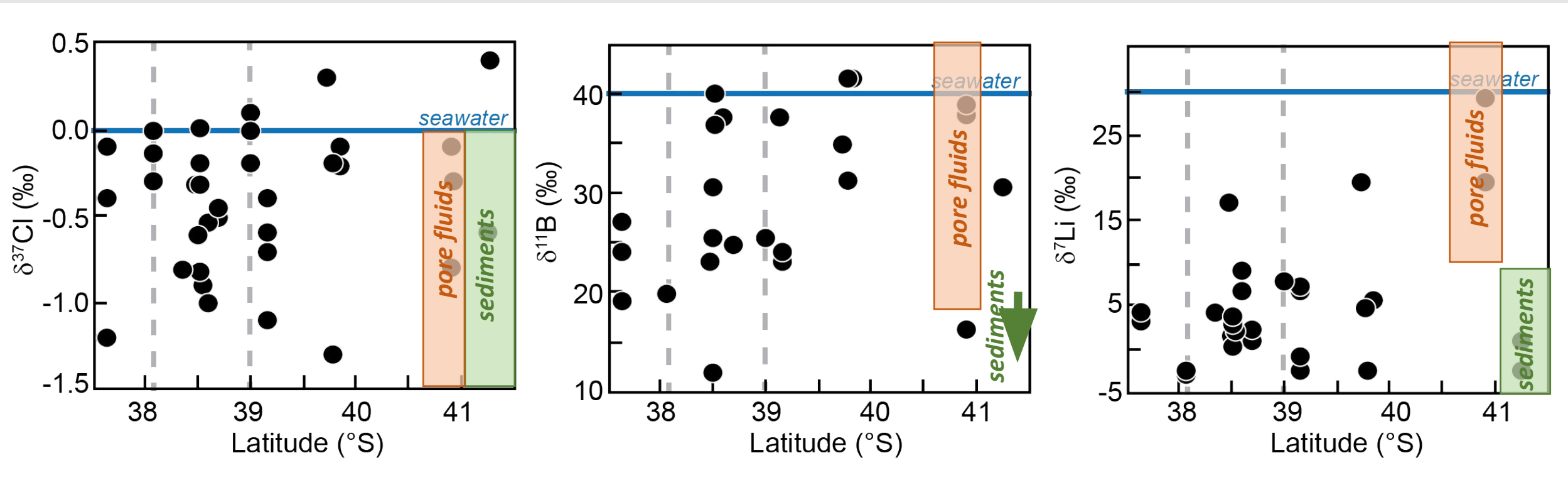


McGuire et al (2017; SZ4D Initiative)



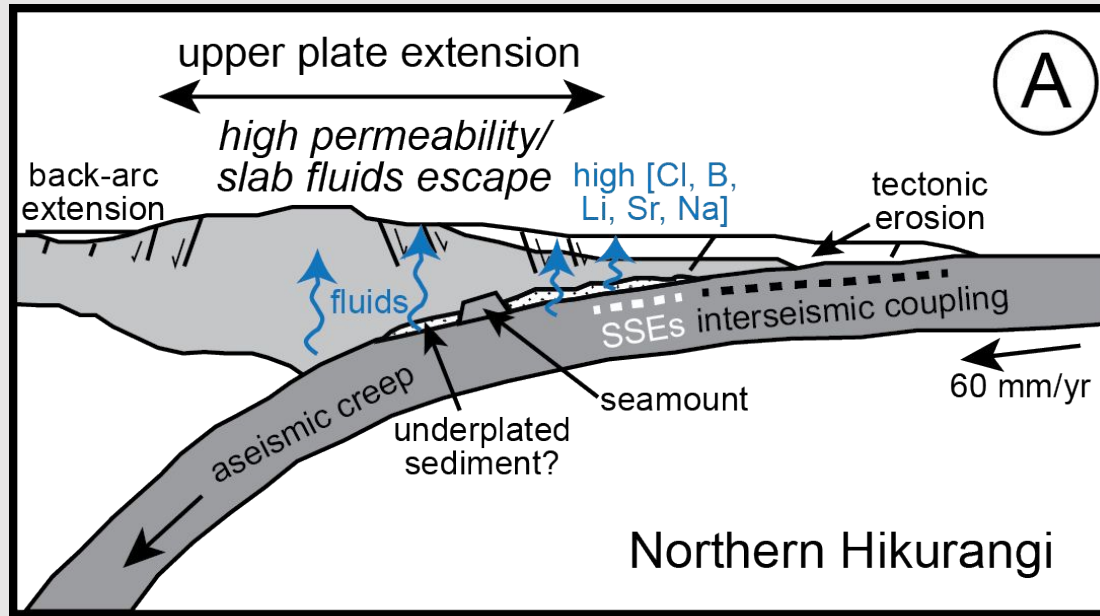


modified from Kendrick et al (2013); additional data from John et al. (2011); Chavrit et al (2016)

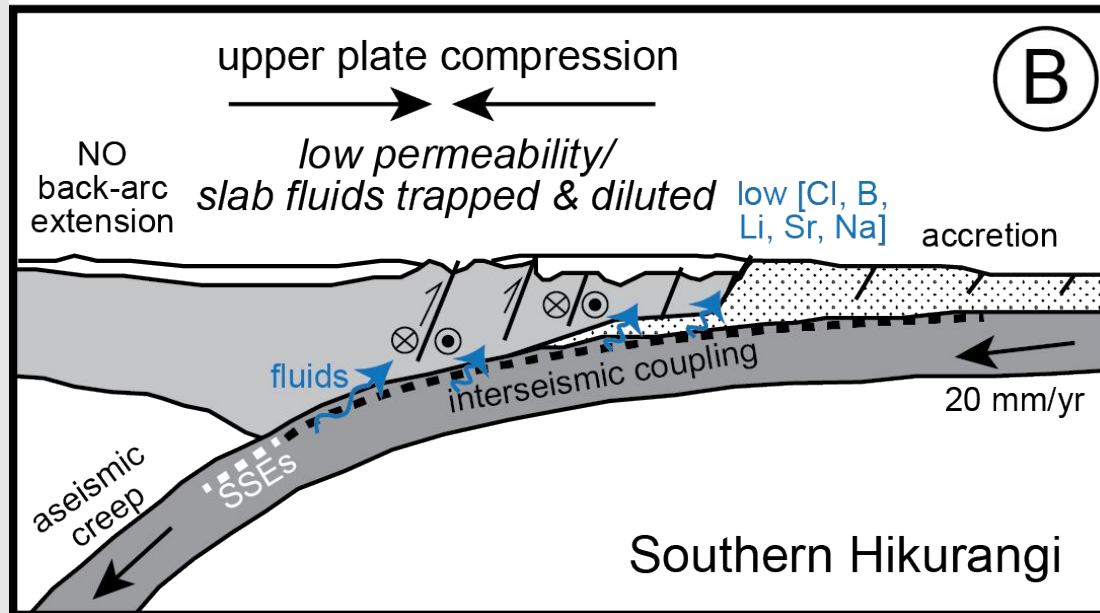


Barnes et al (2019)

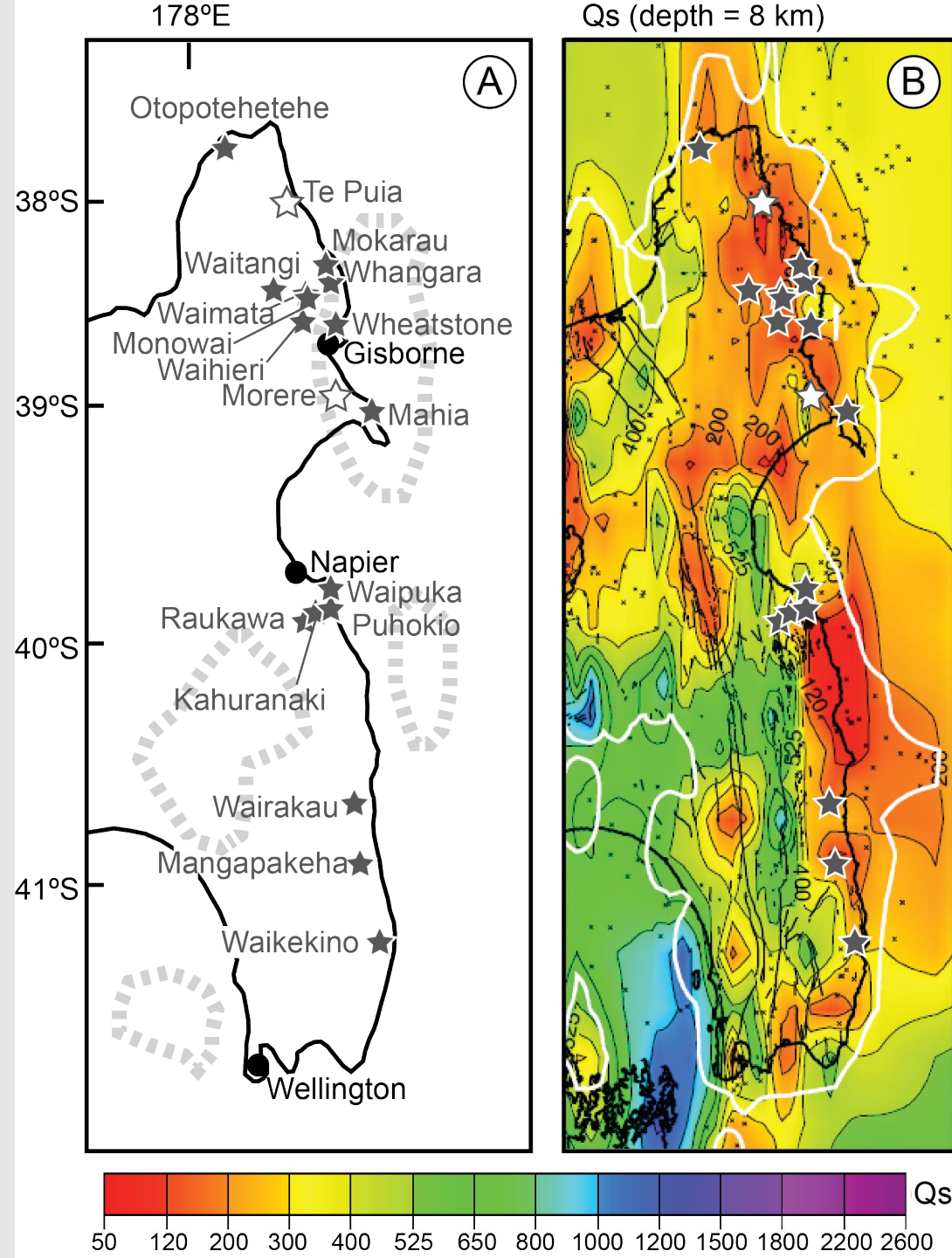
Stable isotope data and elemental ratio data ☐
seawater and sedimentary pore fluids
Not a change in fluid source, but rather fluid flux



Extensional in the North □
 Higher permeability of
 upper plate
 Larger fluid flux



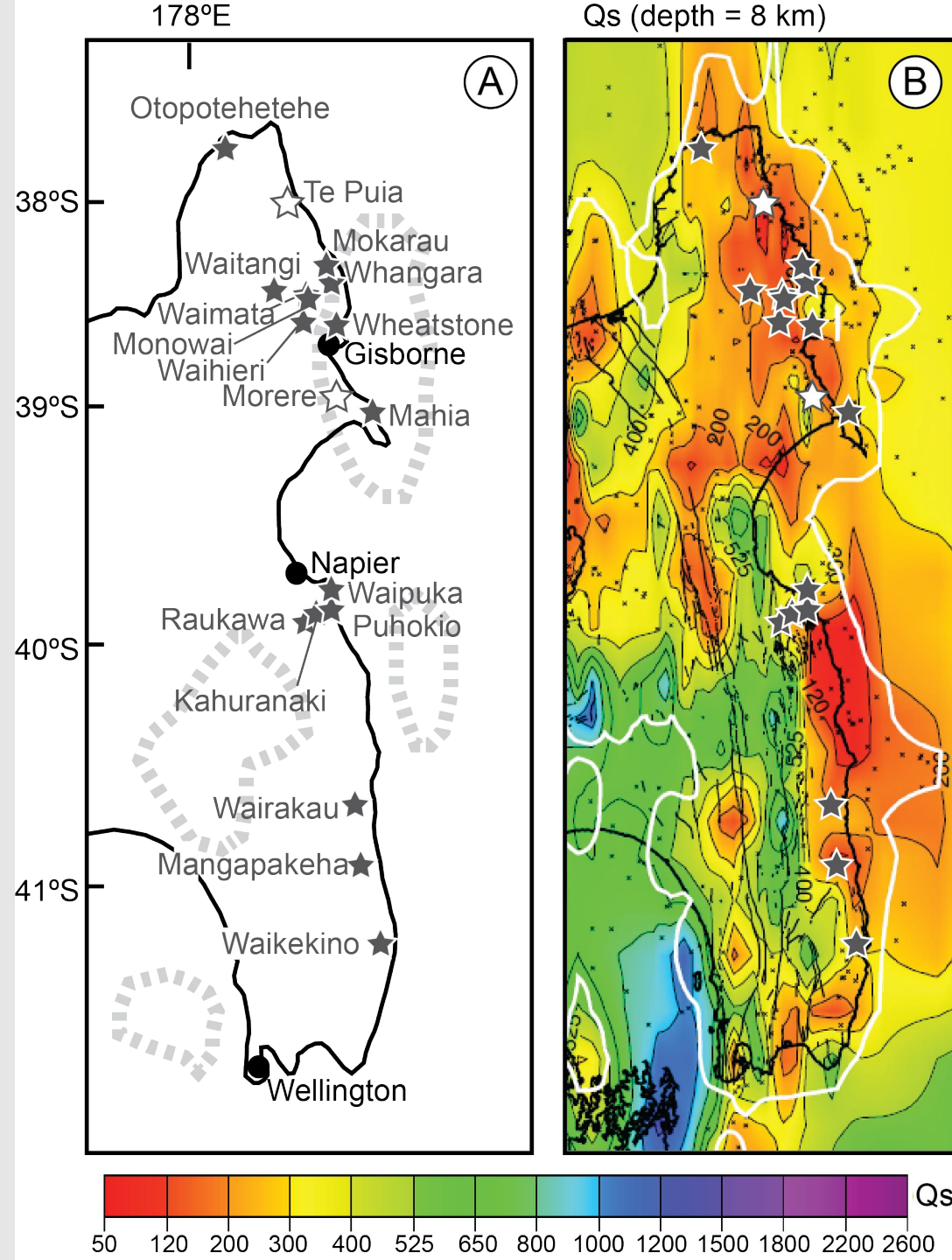
Compressional in the South □
 Lower permeability of upper
 plate
 Lower fluid flux



Low Q (inverse of seismic attenuation) □ high fluid content

Distinct concentration highs correlate with low Q

Eberhart-Phillips et al (2017)



Low Q (inverse of seismic attenuation) □ high fluid content

Distinct concentration highs correlate with low Q

□ integration of geochemical data with other data (e.g., geophysical, mechanics)

Eberhart-Phillips et al (2017)

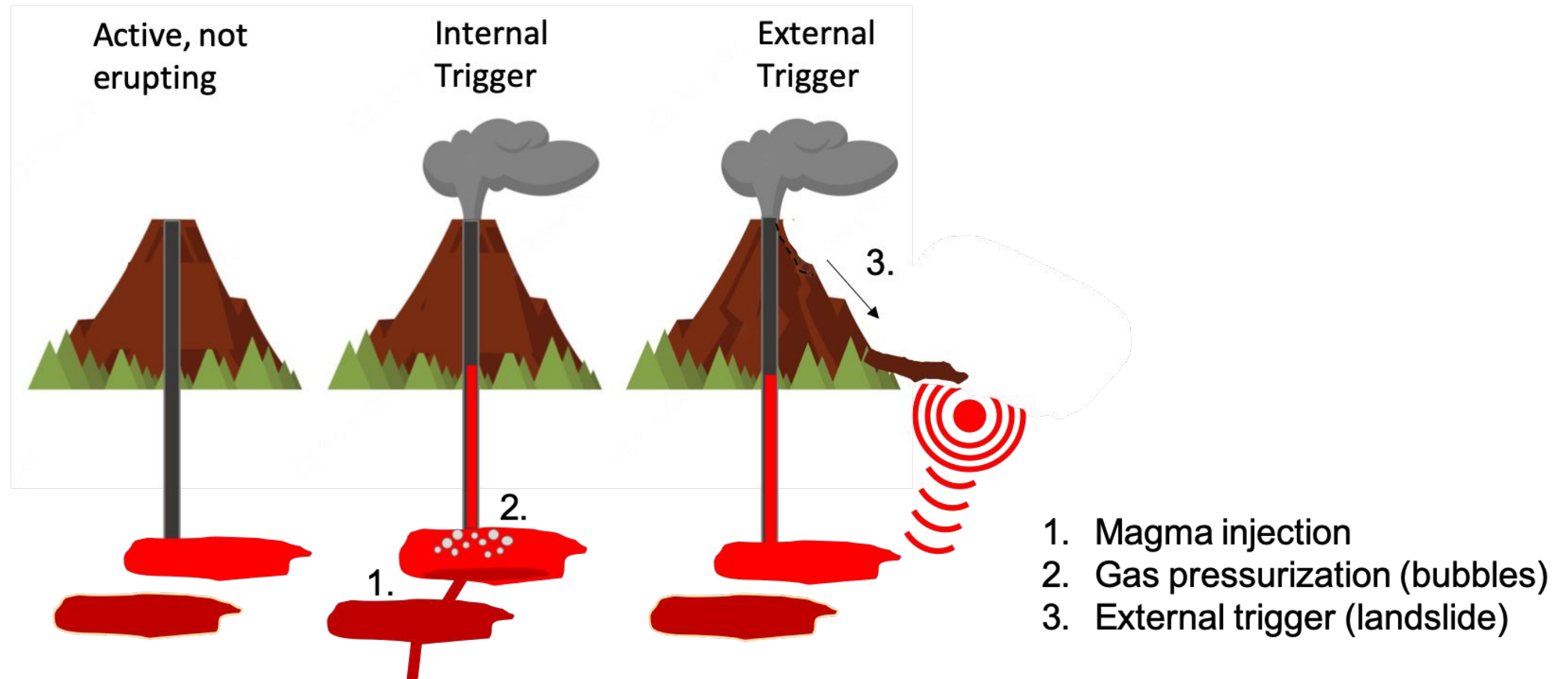
Understanding drivers of geohazards using geochemical methods

(Madison Myers, Montana State University)

Magmas, Melts, and Gases

How do we use the erupted materials from volcanic systems to understand the processes and timescales for triggering volcanic eruptions

Sidebar 3: What drives a volcanic eruption?

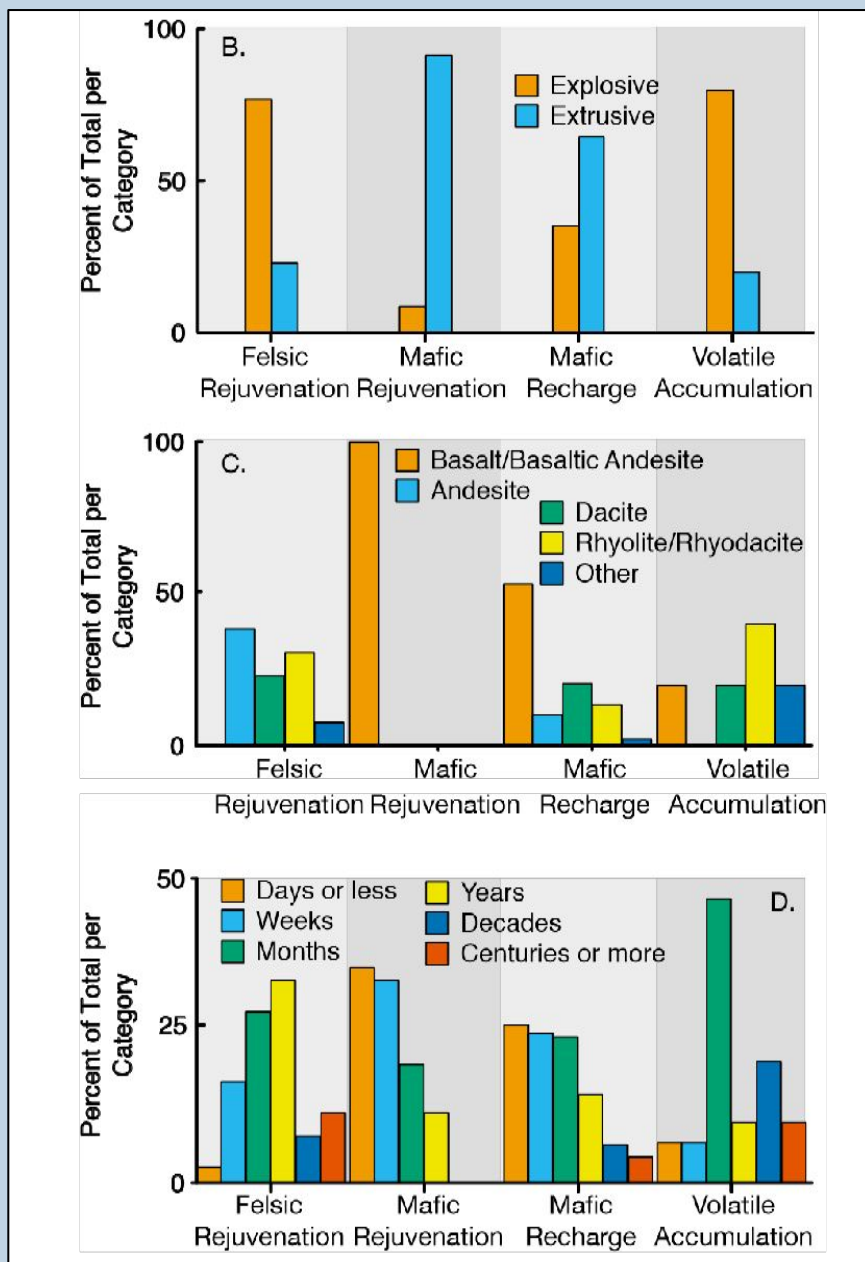


Volcanic eruptions are triggered by both internal processes (from within the magma or magmatic system) and external process (e.g., landslides, earthquake). The most common mechanism for triggering eruptions in arc settings is deeper, hotter magma entering a shallower magma storage region, but eruptions can also be initiated by buildup of gas pressure related to crystallization. One open question volcanologists are trying to constrain is how much time will elapse between the triggering event, and the physical eruption.

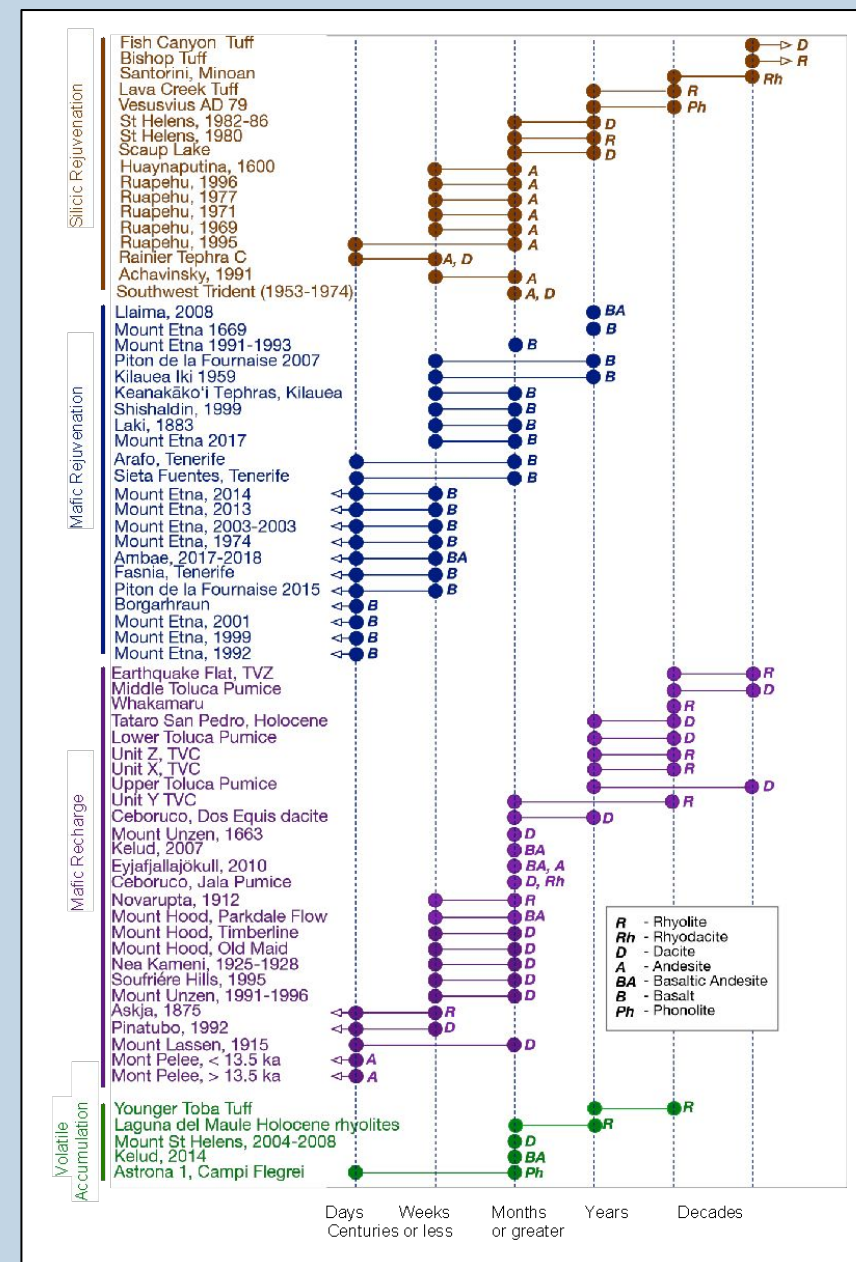
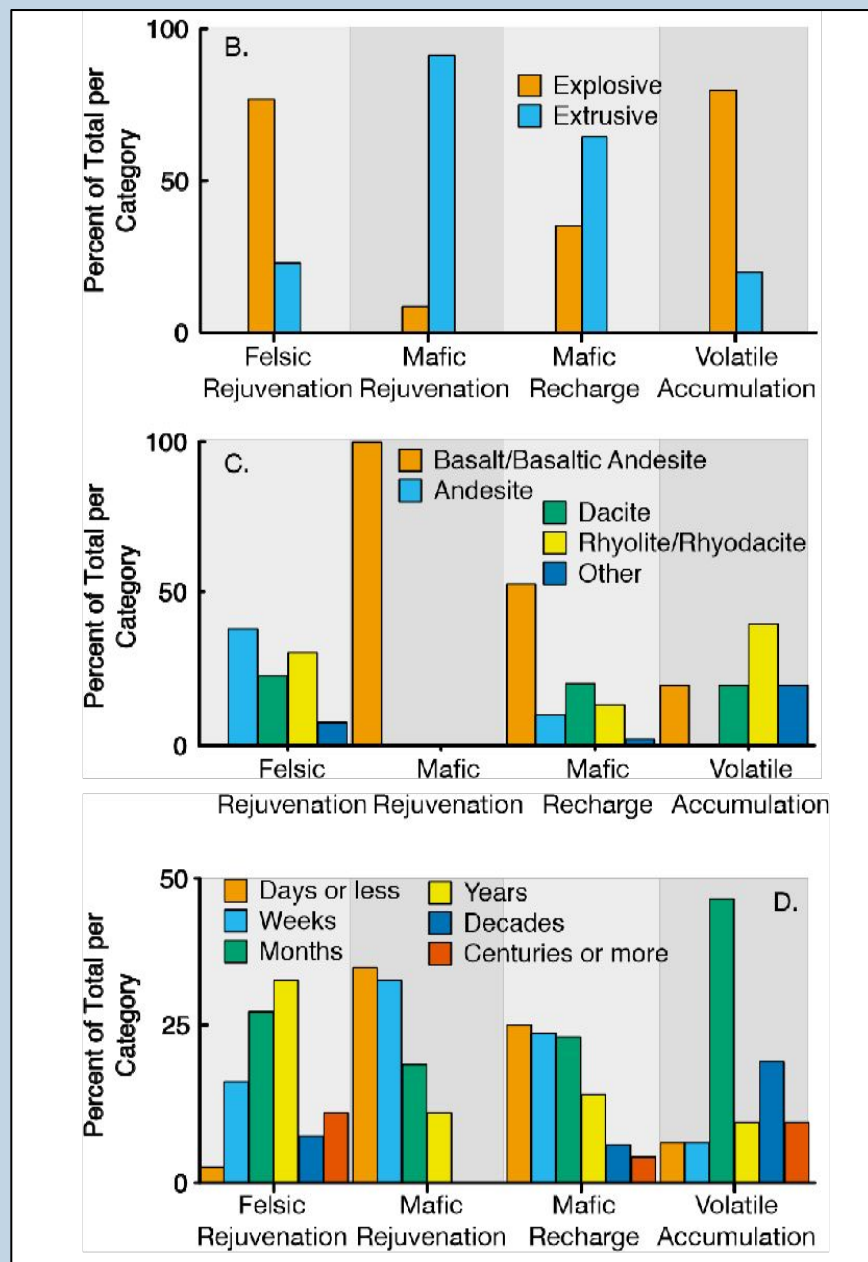
Geochemistry, volatiles and petrology can inform on:

- 1) What triggers and eruption to occur and over what timescale?
- 2) Integrate petrological and geochemical data with geophysics
- 3) Improve our models (diffusion, ascent, overpressurization, etc.) to represent processes
- 4) Measurement, solubility and volatile pathways
- 5) Real time collection and processing

1) What triggers and eruption to occur and over what timescale?

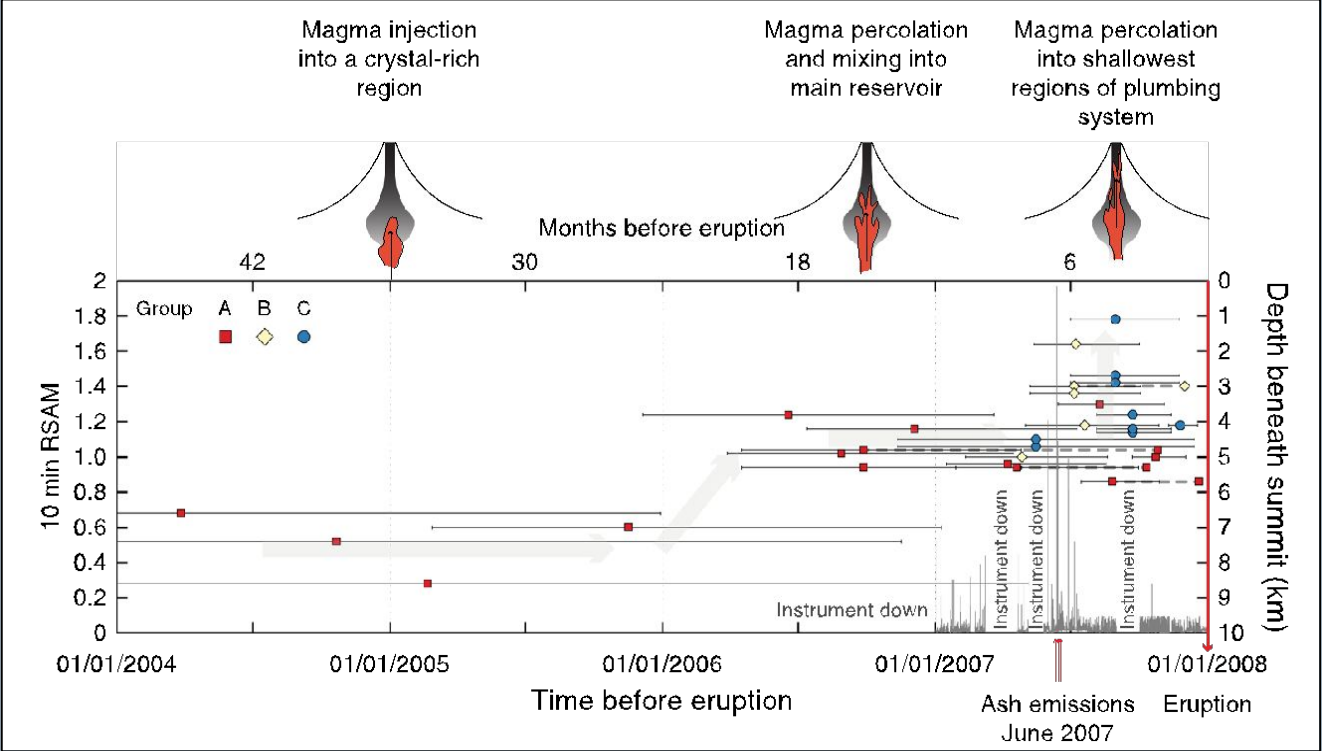


1) What triggers and eruption to occur and over what timescale?



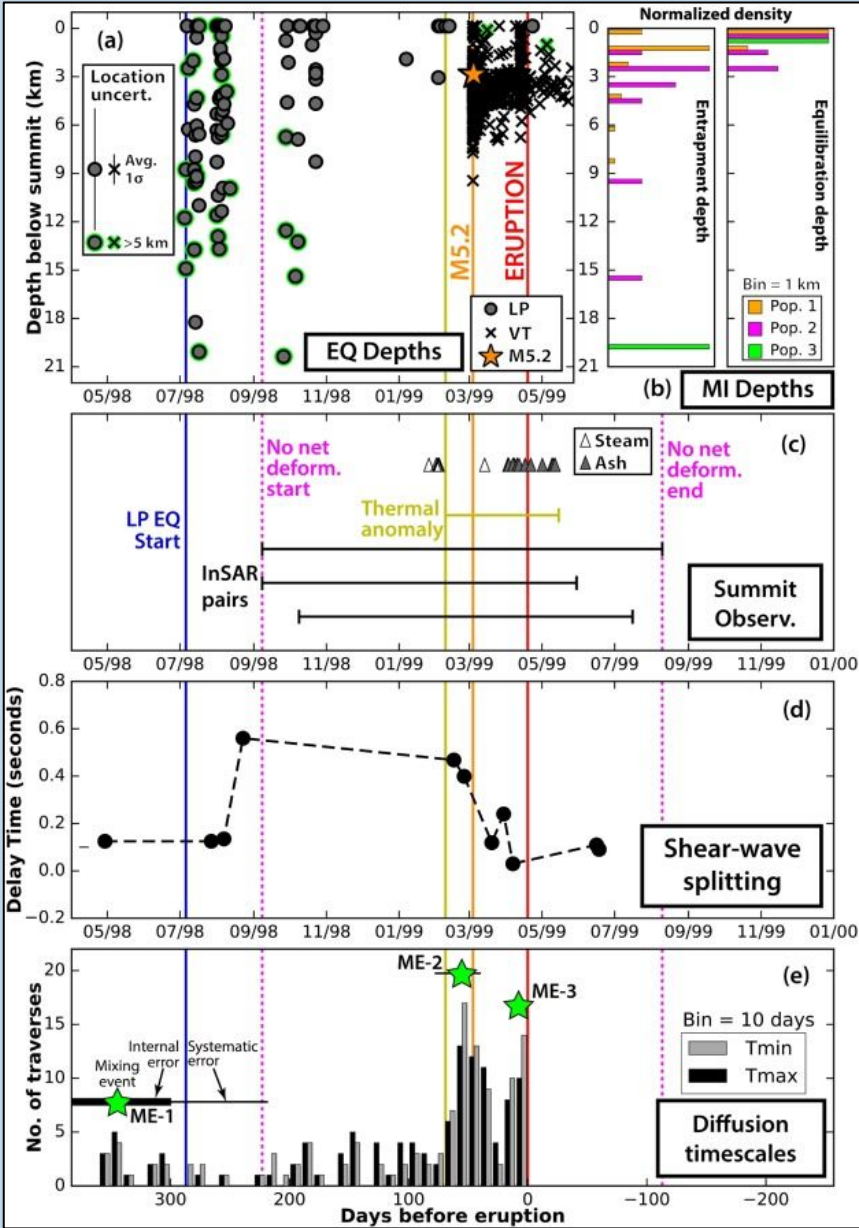
Kent et al. 2020

2) Integrate petrological and geochemical data with geophysics



Ruth et al. 2018

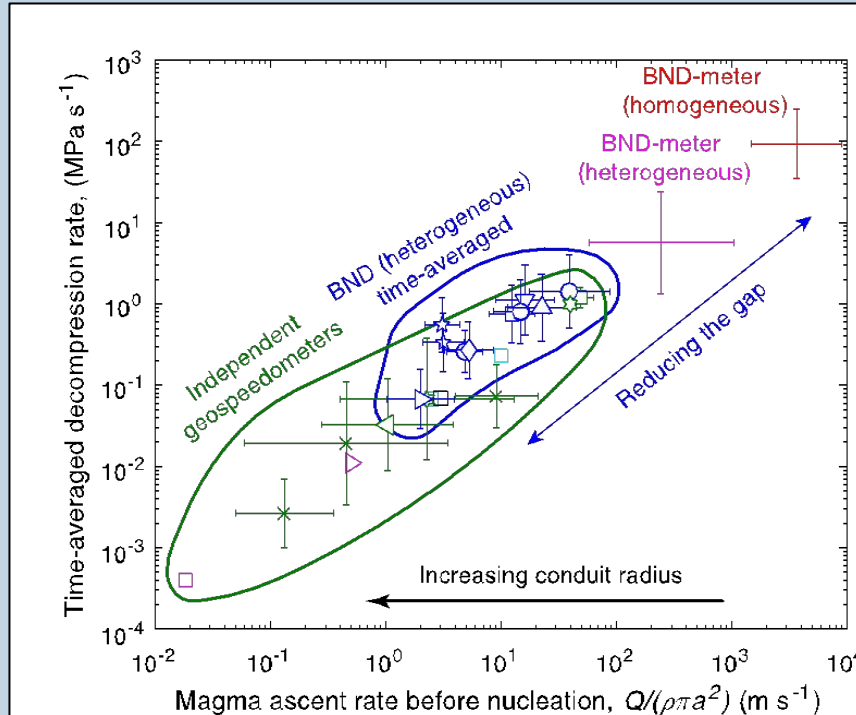
Nature



Rasmussen et al. 2018

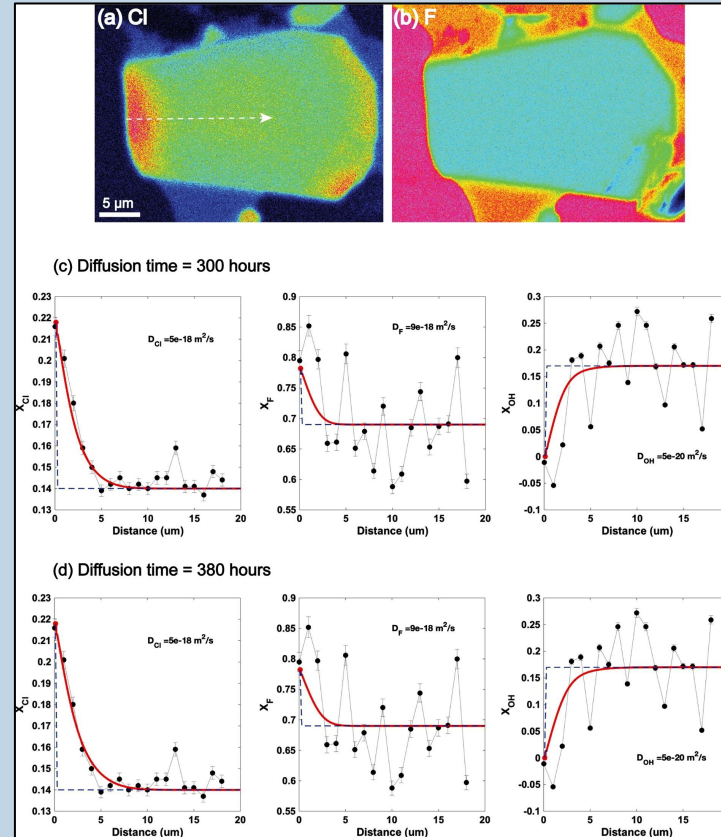
EPSL

3) Improve our models (diffusion, ascent, overpressurization, etc.) to represent processes



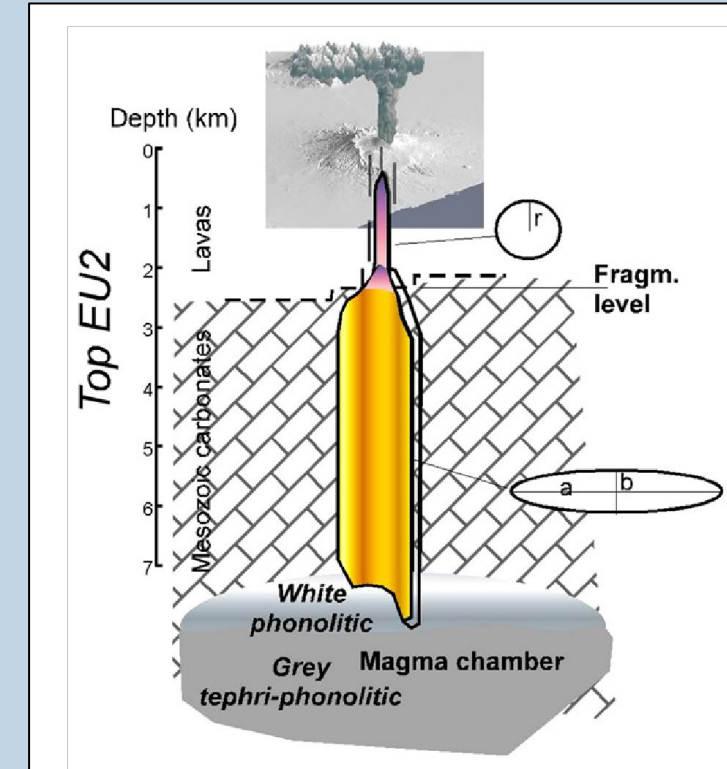
Hajimirza et al. 2021

Nature



Li et al. 2020

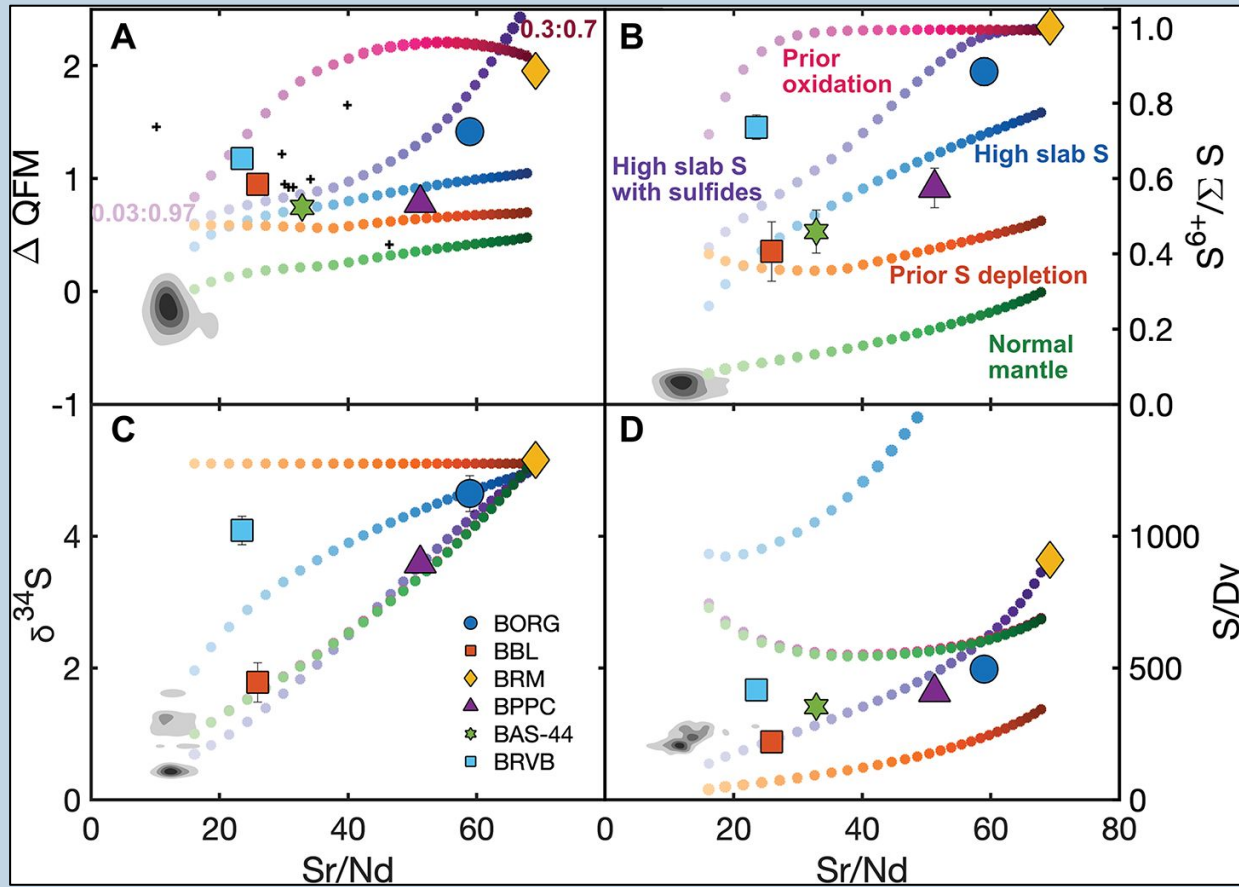
EPSL



Massaro et al. 2018

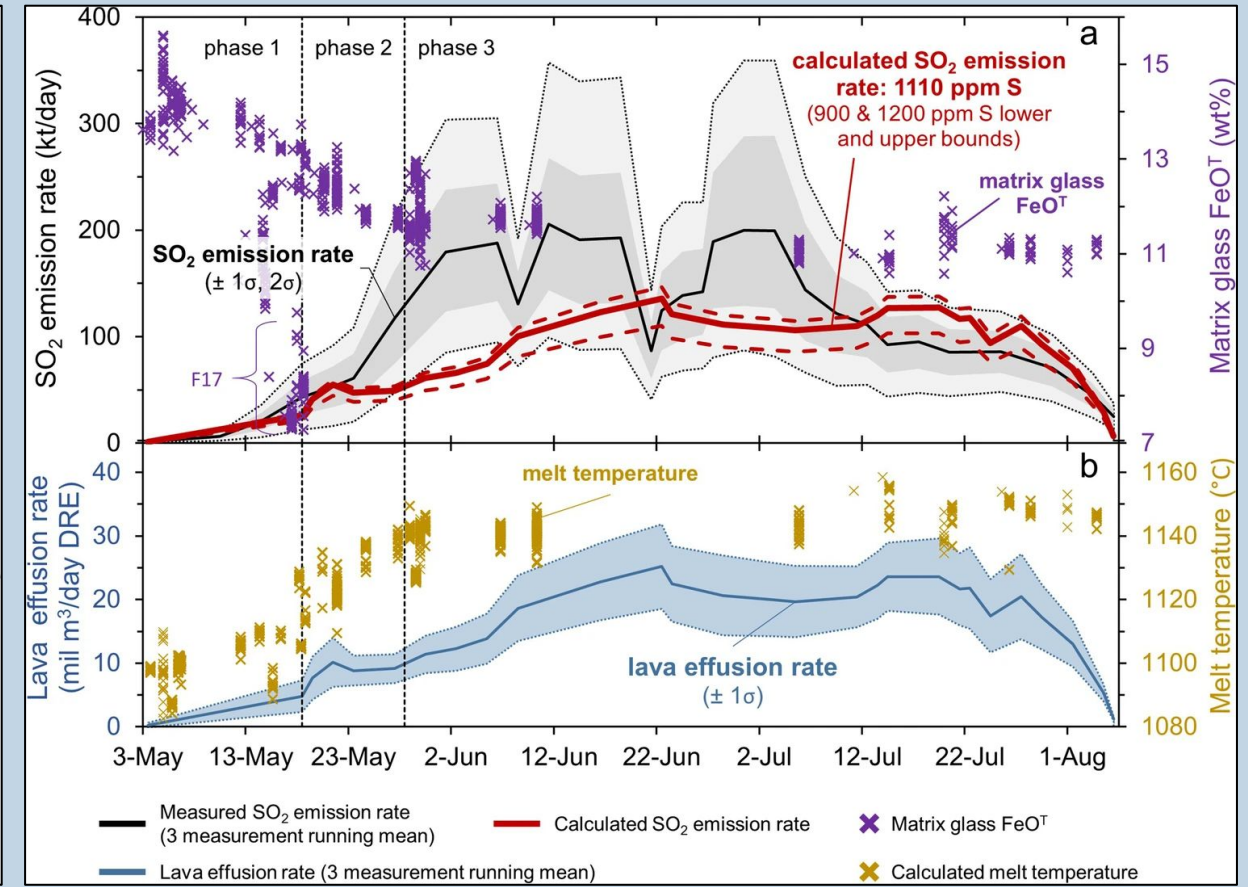
EPSL

4) Measurements, solubility and volatile pathways



Muth and Wallace 2021

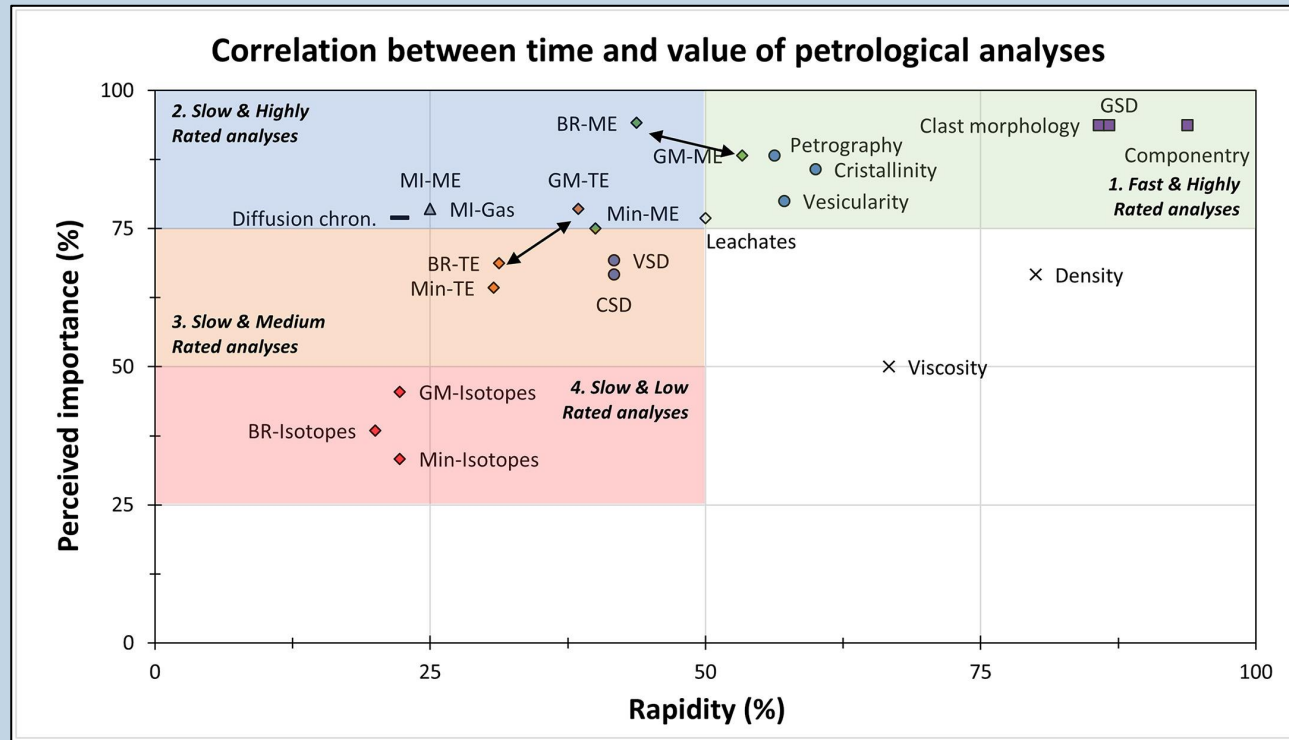
Geology



Lerner et al. 2021

Bull Volc

5) Real time collection and processing

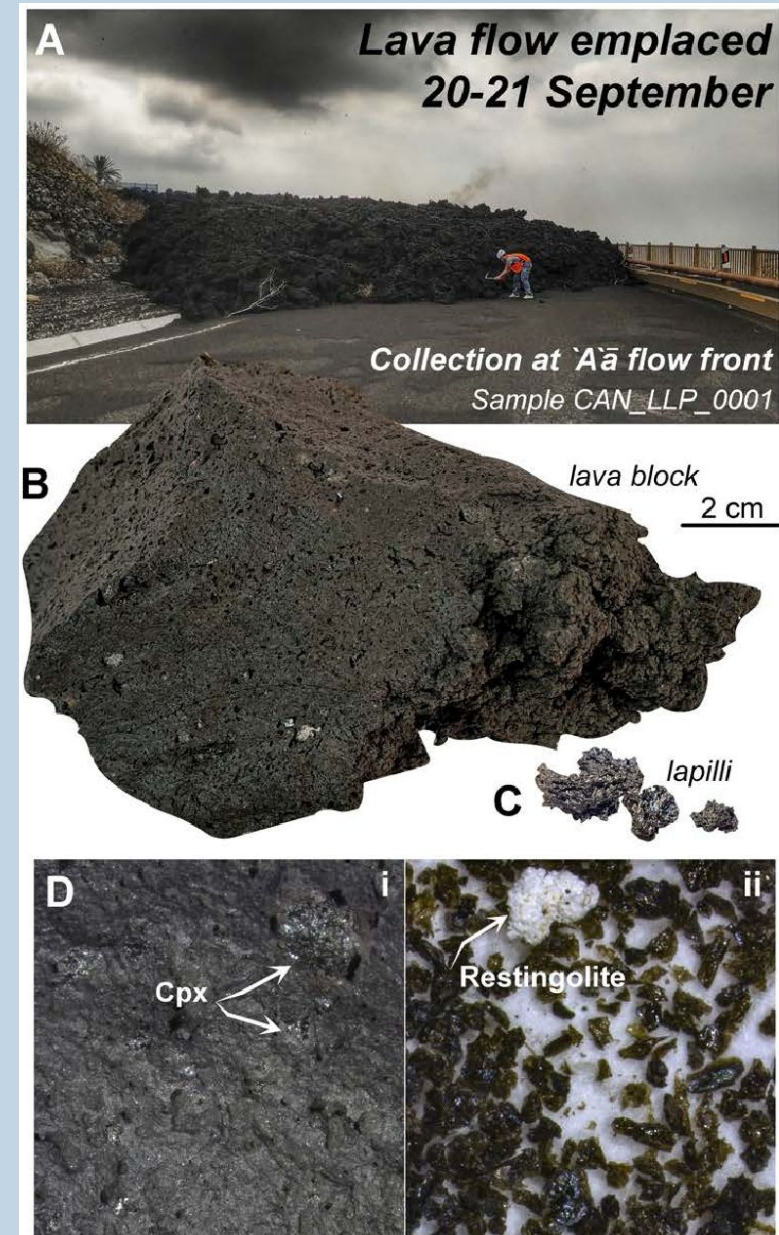


Re et al. 2021

JVGR

Pankhurst et al. 2022

Volcanica



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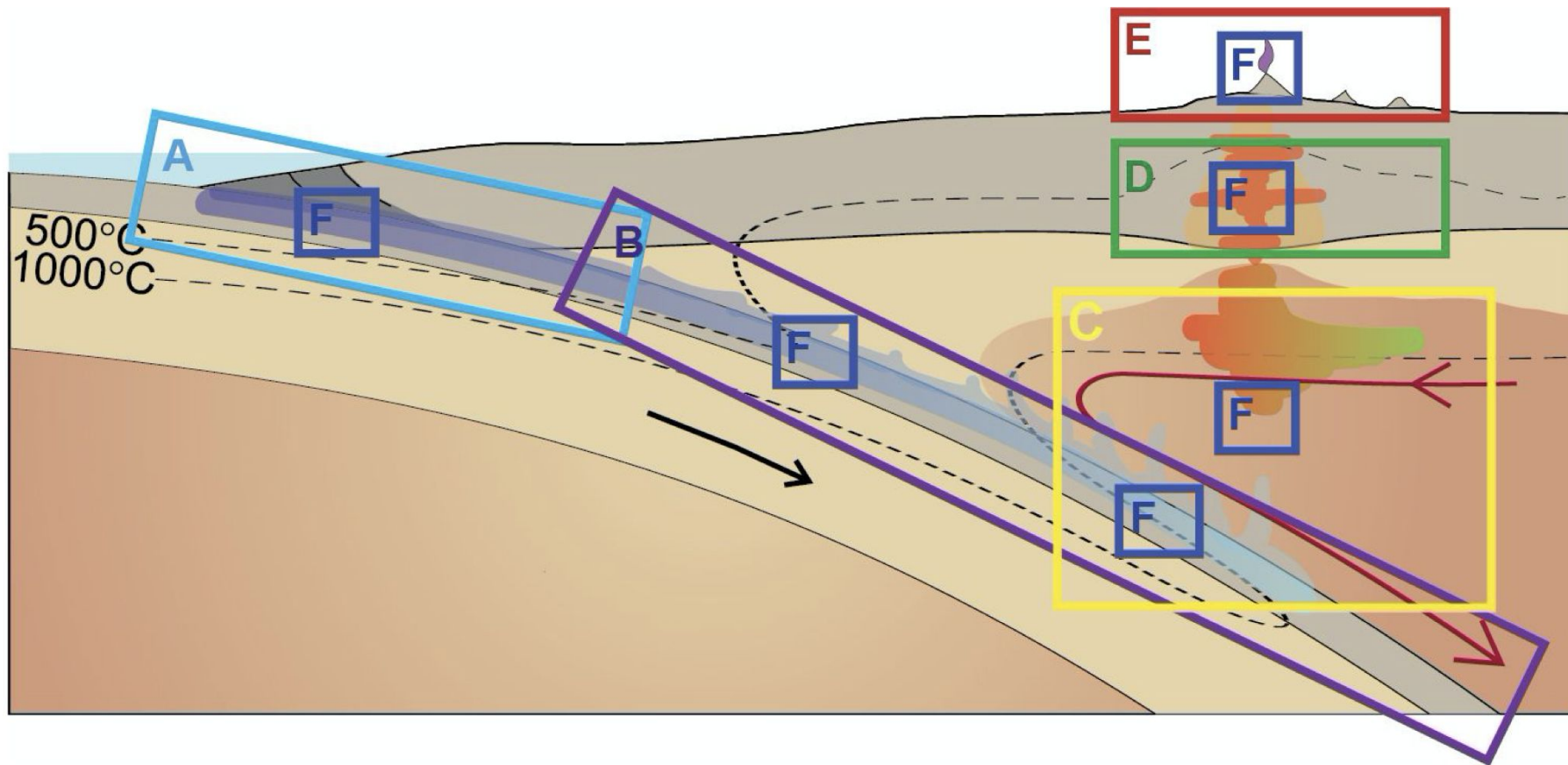


Figure 4. Schematic diagram illustrating the key components of a subduction system and domains (colored boxes with labels A to F) that are commonly investigated as an isolated system of fluid migration. Boxes F indicate example regions where microscopic and/or short-time-scale problems have been investigated.